

Research Article

Magmatic origin and petrogenesis characterization of syenite rock from Pakkanadu alkaline complex, Southern Granulite Terrain, India: Implication on emplacement and petrogenetic history

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Abstract

The present study mainly focused on understanding the magmatic origin and petrogenesis characterization based on the Petrography, major, trace and Rare Earth Element (REE) signatures in the alkaline syenite from Pakkanadu alkaline carbonatite complex. The alkaline plutons from South Indian granulite terrain are intruded along with Archaean epidote-hornblende gneisses. The study area was carbonatite complexes of Tamil Nadu and is characterized by a group of rock associations Carbonatite-Syenite-Pyroxenite - Dunite. From Harker various patterns Pakkanadu alkaline complex syenite showed increasing trends of SiO_2 , Al_2O_3 , $\text{Na}_2\text{O} + \text{K}_2\text{O}$ opposite to decreasing order of CaO , Fe_2O_3 , MgO , TiO_2 , P_2O_5 and MnO trend, suggest fractionation of clinopyroxene, hornblende, sphene, apatite and oxide minerals and feldspar that ruled the fractionation. The concentration of trace elements enriched in Large Ion lithophile elements (LILE) (Ba, Sr, and Rb) elements and High Field Strength Elements (HFSEs) indicated that the dyke intrusion by differentiation of magma from a mantle source. Rare earth element (REE) distribution of Light rare earth element (LREE) enriched and High rare earth element (HREE) depleted pattern show strongly fractionated pattern with moderate Eu anomalies. Plots of tectonic discrimination diagrams of Pakkanadu samples fall in the field of syn-COLG field to the VAG syn- COLG field. For the first time, this type of study was carried out in the study region in a detailed manner. The present study significantly exposed the petrography, petrogenesis and magmatic origin process in the Pakkanadu alkaline carbonatite complex.

Keywords: Alkaline Syenite, Pakkanadu Complex, Southern Granulite Terrane, Tectonic setting

INTRODUCTION

In Southern India, Salem and Dharmapuri districts of Tamil Nadu is a large number of Proterozoic syenite from the Alkaline-Carbonate-Ultramafic Complex of Pakkanadu, emplaced within the Precambrian granulite terranes on North East –South West direction (Gopalakrishnan 1996; Gopalakrishnan 1993; Gopalakrishnan and Ganesam 1992; Gopalakrishnan *et al.*, 2002; Jeyabalan *et al.*, 2015; Srinivas *et al.*, 2011; Gangatharan and Anbarasu 2020 a, b). Tectonically, these alkaline igneous rocks are associated with continental rift valleys or divergent continental margins, oceanic

and continental intraplate settings, and subduction zone magmatism (Fitton and Upton, 1987; Woolley, 1987; Burke *et al.*, 2003, 2008; Leelanandam *et al.*, 2006; Vijaya Kumar *et al.*, 2007; Upadhyay, 2008; Woolley and Kjarsgaard, 2008; Ashwal *et al.*, 2016; Ackerman *et al.*, 2017; Ranjan *et al.*, 2018). Their special geochemical signatures correlated with specific tectonic associations serve as an important tool for understanding mantle and crustal processes and their geodynamic manifestations (Santosh *et al.*, 1989; Natarajan *et al.*, 1994; Kumar *et al.*, 1998; Upadhyay *et al.*, 2006a, b; Upadhyay, 2008; Mukhopadhyay *et al.*, 2011; Renjith *et al.*, 2014; Chakrabarty *et al.*, 2016;

Hippe *et al.*, 2016; Ackerman *et al.*, 2017; Das *et al.*, 2019; Ranjan *et al.*, 2018; Schleicher, H. 2019; Krishnamurthy, 2019; Paul *et al.*, 2020). Geologically Pakkanadu Alkaline Carbonatite complex (PACC) ranging in Neo-Proterozoic age is distributed in a region of 150km² occurring 60 km SW of the Samalpatti pluton. Profound crustal breaks (stitch/shear zones) framed during late Archaean to Proterozoic age have been barged in via Carbonatite Alkaline complexes. The Pakkanadu Alkaline Complex (PACC) is metamorphosed as shown by the development of large scale deformational structures (Miyazaki *et al.*, 1999; Pandit *et al.*, 1998; Schleicher *et al.*, 1998; Pandit *et al.*, 2002; Moller *et al.*, 2001). Lithologies of this complex includes intrusive syenite, and discontinuous bodies of carbonatite associated with pyroxenite rocks. Structurally this alkaline complex has been intruded into migmatite gneisses, Hornblende biotite gneiss and amphibolite. Past workers have depicted the general geology and emplacement history of the Pakkanadu alkaline syenite complex (Gopalakrishnan 1996; Gopalakrishnan 1993; Gopalakrishnan and Ganesam 1992; Gopalakrishnan *et al.*, 2002; Srinivas *et al.*, 2011, Jeyabalan *et al.*, 2015; Gangatharan and Anbarasu 2020 a, b). The present study reports the field occurrence, textural properties and geochemical signatures of alkaline syenite intrusion emplaced into a Pakkanadu alkaline Carbonatite complex. Further, the study demonstrates the major process involved in forming alkaline syenite plutons with the help of textural, major, trace and rare earth elements concentration and structural signatures in this complex.

MATERIALS AND METHODS

Study area

The present study focused on syenite intrusive bodies exposed in and around of Pakkanadu complex of Kamaneri (77°59'20"E-11°47'08"N), 20 km NW of Salem, on the road connecting Omalur – Mecheri, in the Salem district, Tamil Nadu. The syenite pluton occupies about 1 Sq. Km is composed of medium to coarse-grained K-feldspar and plagioclase observed with the help of hand lens. The smaller syenite pluton in Chindamaniyur and Semmandapatty, located 13 km East of Kamaneri pluton, was exposed in the N-S direction (Fig.1). The varying width of pegmatite veins crosses within the syenite body results in the culmination of magmatic activity. Outcrop of pyroxenite was noticed and syenite composition with distinctive porphyritic structure was also noticed across the syenite body. A magmatic flow foliation is available throughout the syenite body, marked and recorded mainly by the shape alignment of K-feldspars (Fig.2).

Analytical methods

A detailed field study was carried out by using a top-sheet (58 E/13) published by the Geological Survey of India. The sampling stations were marked by geographical coordinates using global position system (GPS) (Model: GARMIN 76 CSx). The fresh rock samples were collected from the ideal field exposure for laboratory study. The structural features and field relationship of the alkaline outcrop was studied in the field itself. The collected rock samples were processed for

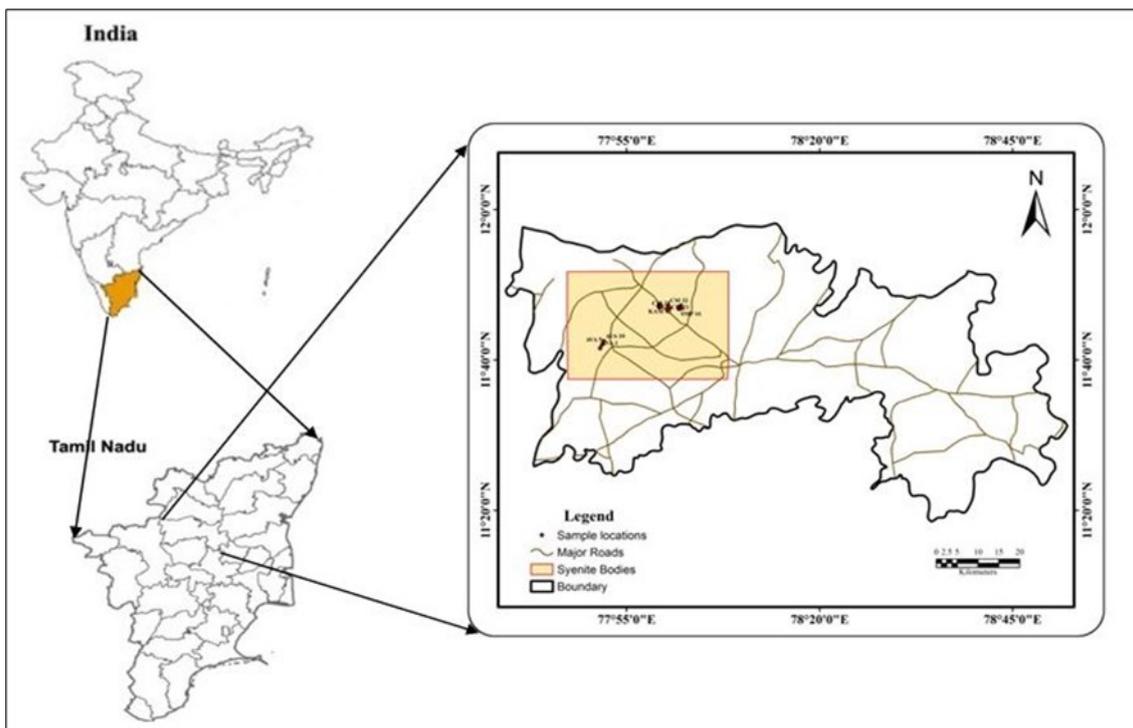


Fig. 1. Location map of the study area

thin section preparation and the petrographic study examined by the petrological microscope LEICA-Model DM 2700P, in Department of Geology, Periyar University, Salem, Tamilnadu. For the petrographic textural study, seventeen samples for whole-rock geochemistry were selected using a standard procedure (Satyanarayanan *et al.*, 2014). Major trace and REE analysis were carried out in National Geophysical Research Institute (NGRI), Hyderabad. The major oxides were analyzed by X-ray fluorescence spectrometer (Philips Magi X PRO mode PW 2440). Trace and REE analysis of whole-rock was done by ICP-MS (Perkin Elmer SCIEX, Model ELAN[®] DRC-II ICP - Mass Spectrometer).

RESULTS AND DISCUSSION

Petrography

The syenite rock from Pakkanadu alkaline complex showed heterogenetic, leucocratic to melanocratic appearance and medium to very coarse-grained pegmatite phase minerals composed mainly of K-feldspar, plagioclase followed by clinopyroxene (aegirine - augite), amphibole with minor accessory of biotite, euhedral titanite, sphene and opaque mineral. The thin section showed that minerals were euhedral to subhedral crystals arranged equally and showed holocrystalline

nature that revealed hypidiomorphic texture. Alkali feldspar showed a wide range of complex exsolution and replacement textures. These minerals were easily identified by their typical polysynthetic and simple twin, and cross-hatched twinning represented and was common in microcline. Some sections of feldspar minerals had inclusions by irregular intergrowths and cryptoperthitic minerals identified in syenite rock intrusion, indicating early-stage exsolution. Feldspar minerals showed intense fractures or cracks indicating solid-state deformation (Fig.3).

Geochemistry:

Major oxides

The major oxides analyses and trace elements of the representative samples from Pakkanadu alkaline complex results are given in the Table (1 and 2). The intrusive rocks from the pluton had a range of silica contents, from approximately 60 to 65 %. The rocks were alkali rich ($\text{Na}_2\text{O} + \text{K}_2\text{O} = 8.66\text{--}13.65\%$) and the plot in the syenite compositional field in the total alkalis versus silica diagram clearly showed that the plots for all investigated syenite rocks fall in the alkaline field as per Middlemost (1994) (Fig. 4). Generally, Syenite rock was clearly pottassic in nature with $\text{K}_2\text{O} > \text{Na}_2\text{O}$. The total alkali ($\text{Na}_2\text{O} + \text{K}_2\text{O}$) vs SiO_2 (Fig. 5) obviously illustrate that the plots for all investigated syenite rocks

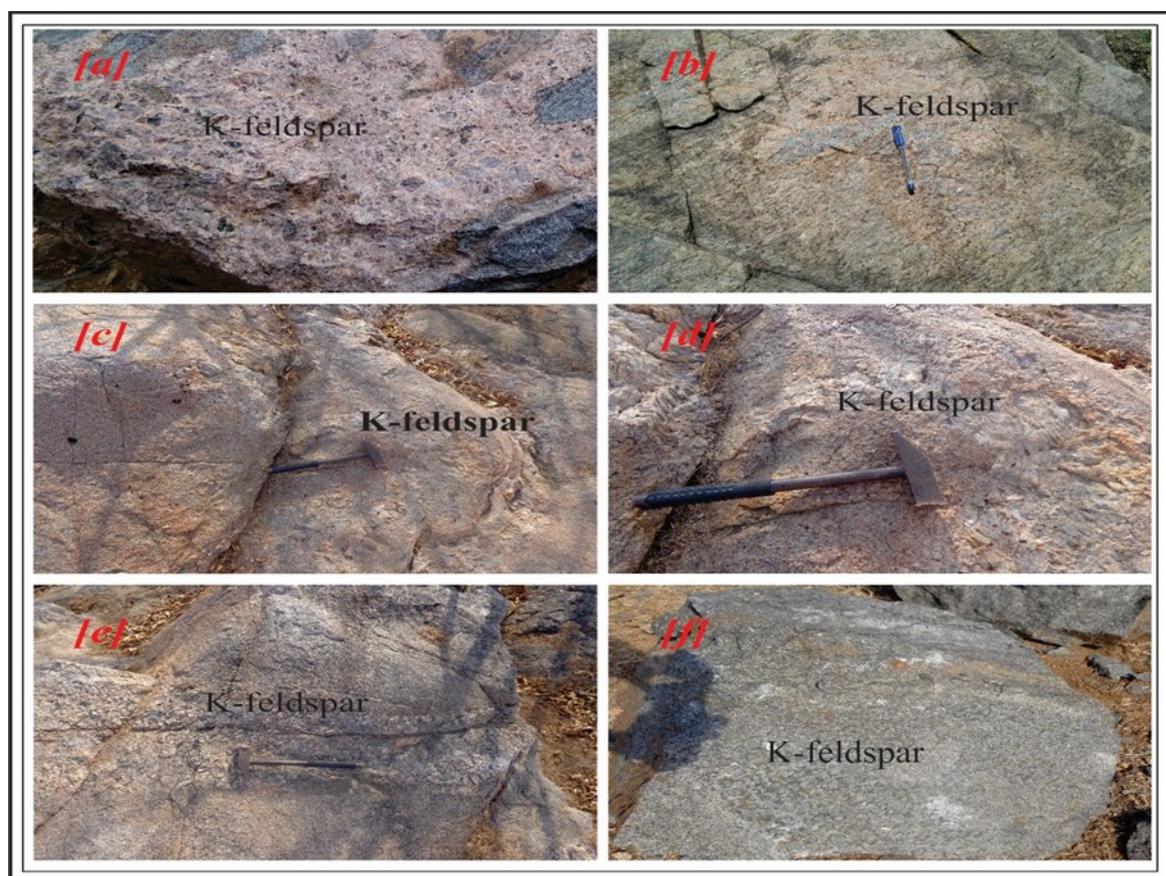


Fig. 2. (a-f). Field photographs of lithology from Pakkanadu alkaline syenite complex

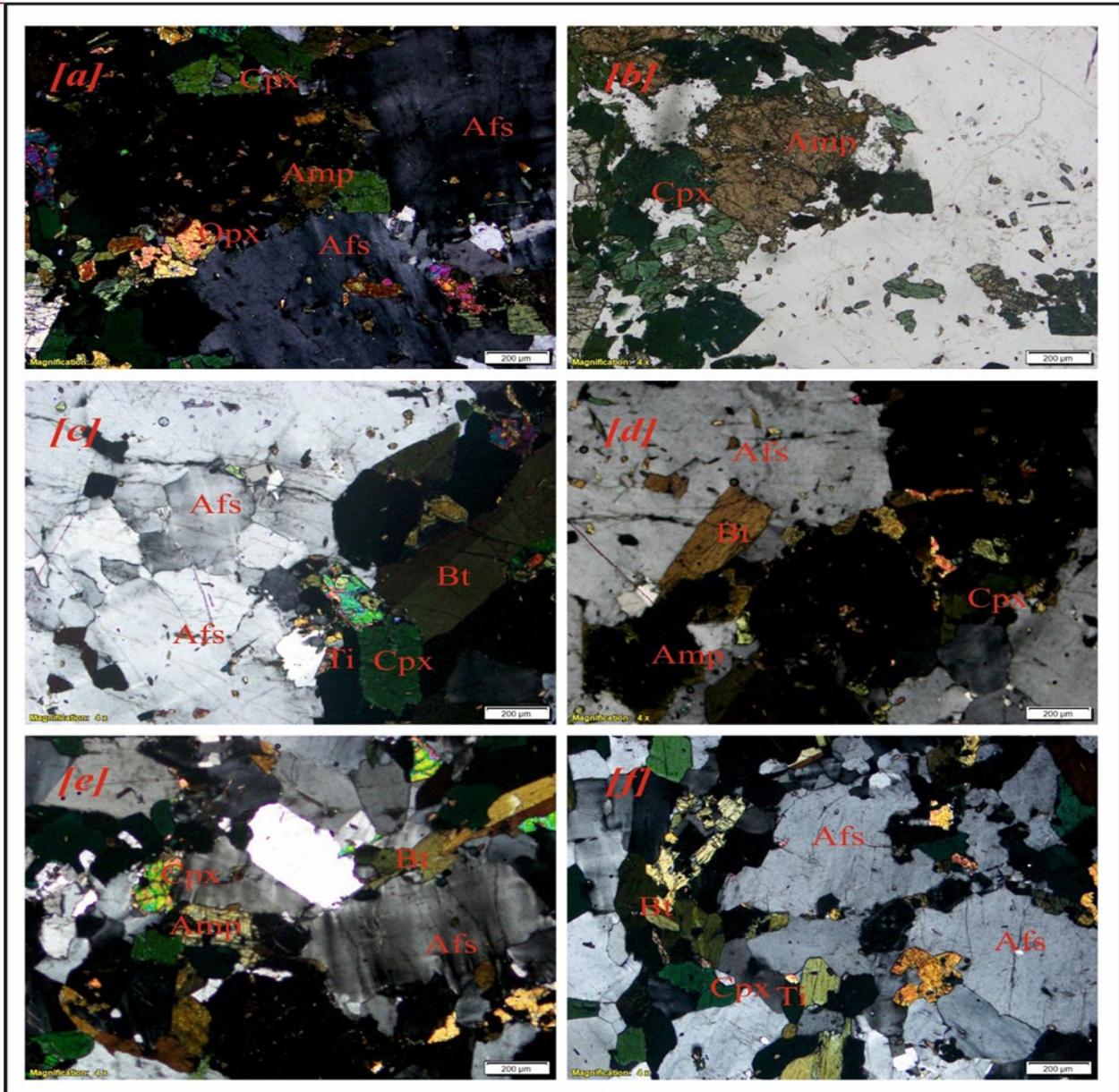


Fig. 3. Photomicrograph of Pakkanadu alkaline syenite a-f shows K-feldspar, plagioclase, clinopyroxene (aegirine - augeite), amphibole and accessory of biotite, euhedral titanite, sphene and opaque mineral are euhedral to subhedral crystals arranged equally shows holocrystalline nature hypidiomorphic texture.

fall in the field of 'alkaline' series nature. The Harker variation diagrams demonstrated the syenite behaviour when they are plots against SiO_2 (Fig. 6). Major oxides concentration showed clear trends of decreasing with increasing against silicate oxides Al_2O_3 , K_2O , CaO , Fe_2O_3 , MgO , Na_2O , TiO_2 , P_2O_5 and MnO contents. These major elements against the SiO_2 trend are generally moderately smooth, implying that the main process involves crystal fractionation results of magmatic evolution. Similar observations were noticed in log $\text{CaO}/(\text{Na}_2\text{O}+\text{K}_2\text{O})$ vs. SiO_2 diagram (Fig.7) wherein all the samples were clustered with alkaline series, which indicates that the study area lithological composition closely belongs to alkaline nature. In the Alkali Iron

Magnesium (AFM) ternary diagram (Fig.8), all the samples clustered with Calc- Alkaline series indicate that the Pakkanadu alkaline lithological complex belongs to plutonic origin and rocks are formed by the slow cooling process during liquid to solid state (Rollinson 1993; Rao and Narayana 2002; Miyazaki *et al.*, 2003; Miyazaki and Santosh 2005; Paul *et al.*, 2020).

Trace elements

Trace elements distribution in primitive mantle normalised spider diagram as proposed by Kerrich and Wyman (1997) The samples are enriched in large ion lithophile elements (LILEs) such as Sr, Ba and Rb, High Field Strength Elements (HFSEs) such as Y, Nb, Zr,

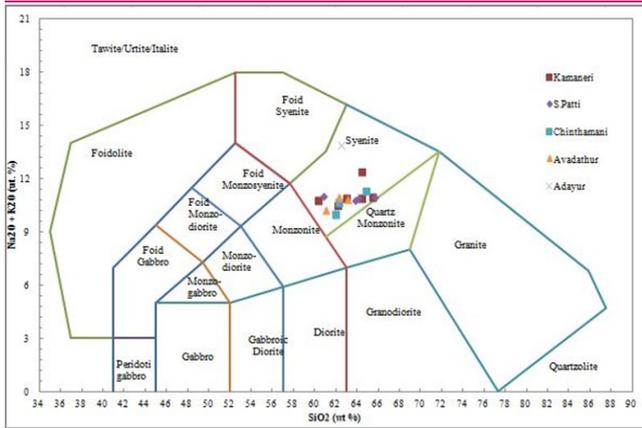


Fig. 4. Diagram of (TAS) Total Alkali vs Silica (after Mid-dlemost, 1994)

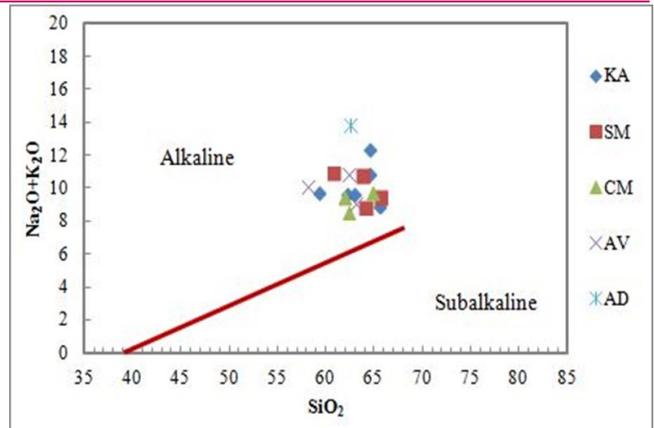


Fig. 5. Plots of ($Na_2O + K_2O$) vs SiO_2 (after Irvine and Baragar 1971).

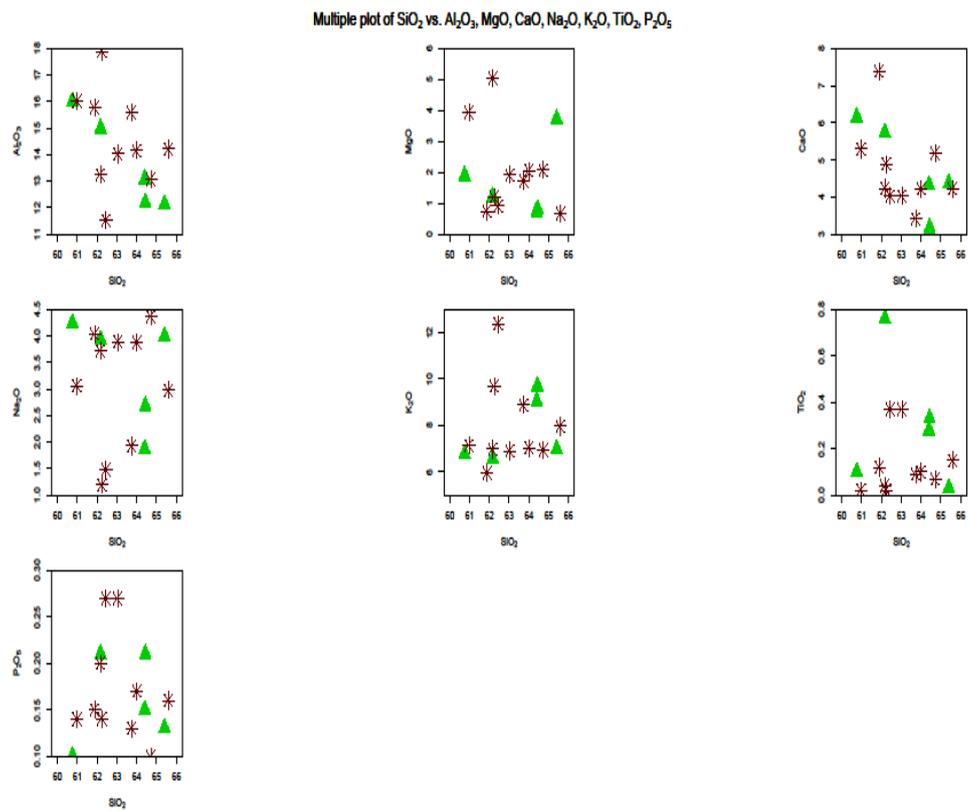


Fig.6. Harker variation diagram of Pakkanadu alkaline Kamaneri Syenite complex

Ta, Hf and, Th, and depletion of Ferromagnesium and Transitional Elements (FTEs) such as V, Cr, Co, Ni, Cu, Zn except Pb (Fig.9). The high concentration of enriched LILE (Ba, Sr, and Rb) elements indicates that the dyke intrusion is derived from magma differentiation from mantle sources. The high Ba and Sr anomaly concentration infer that lithological characterization of Pakkanadu alkaline complex is mainly controlled by plagioclase fractionation and followed by a small range of partial melting. The depletion of ferromagnesian and transitional elements such as V, Cr, Co, Ni, Cu, and Zn from the primitive mantle normalized values reveal that mafic minerals deficient, feldspar ruled fractionation as

also reported earlier (Taylor *et al.*, 1967; Rollinson 1993; Schleicher *et al.*, 1998; Pandit *et al.*, 2002; Upadhyay *et al.*, 2006c; Srinivas *et al.*, 2011; Santosh *et al.*, 2014; Paul *et al.*, 2020).

Rare earth elements

The Rare earth element concentrations of the syenite rock are demonstrated in chondrite-normalized diagrams using the chondrite values of McDonough and Sun (1995) (Fig. 10). The Rare Earth Element pattern showed overall low to moderate of the Σ REE contents from alkaline syenite rock values ranges from 25.77-388.49 mean value 160.94 (Table 3). Σ LREE values

range from 22.20-35.5, mean value 145.70, Σ HREE values range from 3.20-36.03, mean value 15.36. The present record on the REE of the studied syenite indicated that all samples are strongly enriched in LREE ($(La/Sm)_n$ (values ranged from 2.11-12.02, mean value 6.74), and depleted in HREE ($(Tb/Yb)_n$ (values ranged from 0.15 - 0.49, mean value 0.33), for Pakkanadu complex alkaline syenite, respectively. This suggests

that garnet was a residual phase during partial melting in the asthenosphere. The chondrite-normalized REE distribution patterns illustrate that the syenite rock samples are characterized by enrichment in the LREE values ranges from 2.11-12.02 mean value 6.87 and depletion in HREE values range from 0.15-0.49 mean value 0.34. The LREE/HREE values range from 5.15-16.30, mean value of 9.80. The Europium (Eu/Eu^*)

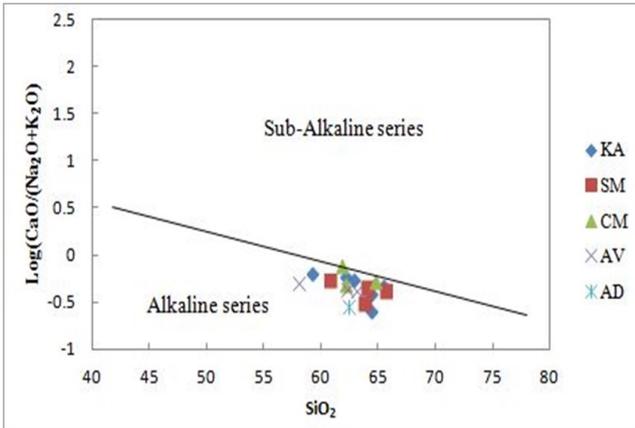


Fig.7. Log (CaO / Na₂O + K₂O) vs SiO₂ of Syenite rock

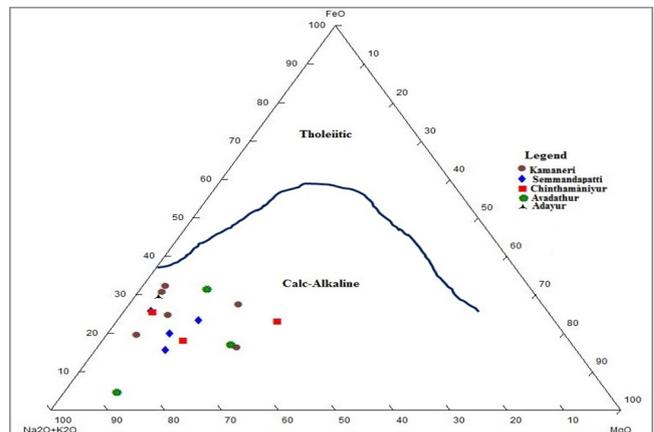


Fig. 8. AFM Trilinear plot (Irvine and Baragar 1971).

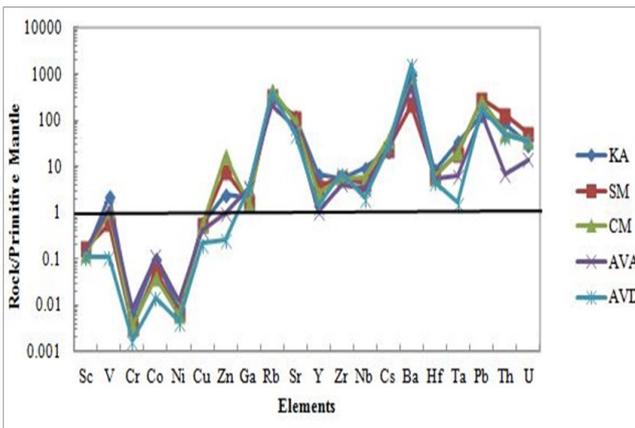


Fig. 9. Primordial Mantle normalized multielement spidergram of Pakkanadu alkaline Kamaneri Syenite complex (Kerrich and Wymann, 1997).

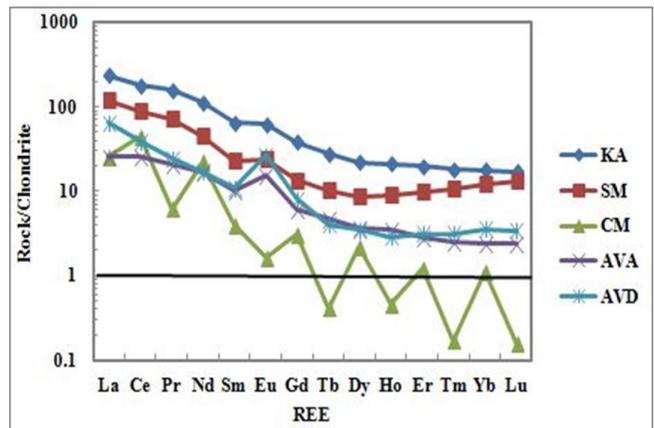


Fig. 10. Chondrite normalized REE patterns of Pakkanadu alkaline Kamaneri Syenite complex (McDonough and Sun, 1995).

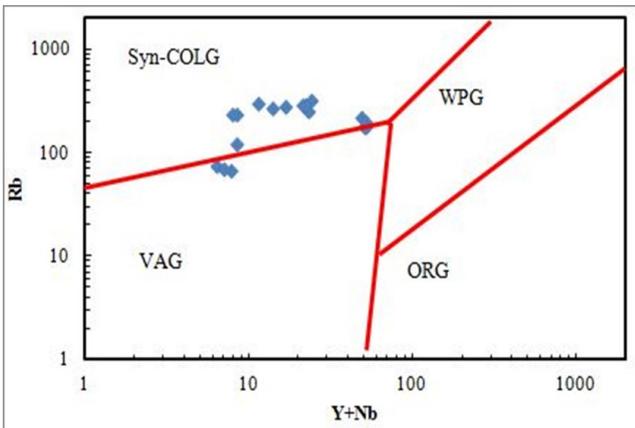


Fig. 11. Tectonic discrimination diagrams of the Pakkanadu alkaline Kamaneri Syenite complex (after Pearce, 1996).

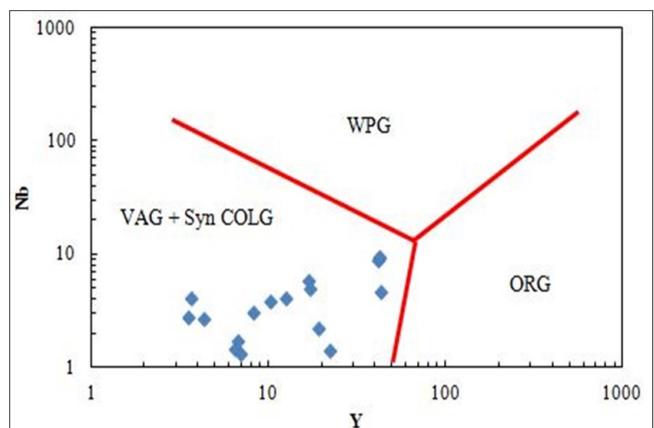


Fig. 12. Nb vs Y plots for tectonic discrimination of Pakkanadu alkaline Kamaneri Syenite complex (after Pearce, 1984).

Table 1. Whole-rock chemical composition of Pakkanadu alkaline complex (W %) at TamilNadu

| Sample | KAM-26 | KAM-31 | KA-35 | KA-30 | SM-1 | SMP-19 | SMP-16 | SMP-35 | CM-21 | CM-25 | CM-32 | AV-10 | AVA-1 | AVA-5 | ADA-1 |
|--------------------------------|--------|--------|-------|-------|-------|--------|--------|--------|-------|-------|-------|-------|-------|-------|-------|
| SiO ₂ | 65.40 | 62.17 | 64.40 | 64.43 | 60.75 | 63.74 | 65.60 | 64.00 | 64.72 | 62.17 | 61.88 | 60.98 | 63.05 | 62.25 | 62.45 |
| Al ₂ O ₃ | 12.12 | 14.98 | 13.07 | 12.19 | 16.00 | 15.59 | 14.21 | 14.15 | 13.06 | 13.23 | 15.77 | 16.02 | 14.02 | 17.88 | 11.52 |
| Fe ₂ O ₃ | 2.43 | 3.55 | 5.51 | 5.81 | 2.37 | 3.08 | 3.51 | 3.31 | 2.60 | 4.04 | 3.48 | 2.85 | 5.02 | 0.57 | 6.11 |
| MnO | 0.03 | 0.04 | 0.07 | 0.07 | 0.02 | 0.04 | 0.04 | 0.05 | 0.04 | 0.06 | 0.04 | 0.03 | 0.04 | 0.01 | 0.08 |
| MgO | 3.72 | 1.19 | 0.71 | 0.81 | 1.89 | 1.72 | 0.66 | 2.05 | 2.08 | 5.04 | 0.74 | 3.93 | 1.91 | 1.20 | 0.91 |
| CaO | 4.36 | 5.73 | 4.32 | 3.17 | 6.15 | 3.41 | 4.22 | 4.22 | 5.18 | 4.23 | 7.38 | 5.29 | 4.03 | 4.88 | 4.03 |
| Na ₂ O | 3.99 | 3.93 | 1.87 | 2.68 | 4.24 | 1.93 | 2.98 | 3.88 | 4.36 | 3.72 | 4.04 | 3.06 | 3.89 | 1.20 | 1.48 |
| K ₂ O | 6.98 | 6.56 | 9.04 | 9.70 | 6.77 | 8.89 | 7.97 | 7.04 | 6.94 | 6.98 | 5.96 | 7.13 | 6.90 | 9.70 | 12.37 |
| TiO ₂ | 0.03 | 0.76 | 0.28 | 0.33 | 0.10 | 0.09 | 0.15 | 0.10 | 0.07 | 0.04 | 0.12 | 0.02 | 0.37 | 0.02 | 0.37 |
| P ₂ O ₅ | 0.13 | 0.21 | 0.15 | 0.21 | 0.10 | 0.13 | 0.16 | 0.17 | 0.10 | 0.20 | 0.15 | 0.14 | 0.27 | 0.14 | 0.27 |
| LOI | 0.30 | 0.19 | 0.57 | 0.48 | 0.57 | 0.36 | 0.43 | 0.30 | 0.34 | 0.18 | 0.39 | 0.50 | 0.39 | 0.32 | 0.39 |
| Total | 99.49 | 99.31 | 99.96 | 99.87 | 98.93 | 98.99 | 99.93 | 99.27 | 99.49 | 99.88 | 99.96 | 99.95 | 99.89 | 98.15 | 99.98 |

anomaly of the study area alkaline syenite rocks had limits of concentration enriched with values ranging from 0.18-0.50 mean value 0.26. This can possibly be related to mixing with plagioclase-rich melt. This may be related to the later formation of this intrusive body and strong fractionation during crystallization. From the REE spider diagram, all the REE elements observed in the steep slope pattern, all the samples observed higher concentration of LREE and depleted concentration of HREE. This enriched and depleted pattern indicated that alkaline magma was derived from the mantle sources as also reported by earlier workers (Taylor *et al.*, 1967; Rollinson 1993; Schleicher *et al.*, 1998; Pandit *et al.*, 2002; Upadhyay *et al.*, 2006c; Srinivas *et al.*, 2011; Santosh *et al.*, 2014; Paul *et al.*, 2020).

Tectonic setting

The emplacement tectonics of various igneous suits are demonstrated by different bivariant and triangular diagrams with the help of major oxides and trace elements (eg. Wood *et al.*, 1979; Floyd and Winchester, 1975; Pearce and Cann 1973; Pearce *et al.*, 1977, 1984; Pearce, 2008; Vrublevskii *et al.*, 2019). In the global context, the syenite magmatism corresponding to Archean age were quite rare compared to Proterozoic, whereas the latter was commonly associated with the continental rift/extensional tectonic settings. The present study used the tectonic discrimination diagrams (Pearce *et al.*, 1984 and 1996) to identify the tectonic emplacement. All the samples were plotted (Y+Nb) vs Rb (Fig.11) and Y vs Nb (Fig.12), most of the samples clustered in the field of syn-COLG + VAG and VAG + syn-COLG, respectively. These samples association of syn-COLG + VAG and VAG + syn-COLG confirm that the Pakkanadu alkaline carbonate complex was mostly derived from fractionate of mantle-derived magmatic sources.

Conclusion

The field characterization of exposure from Pakkanadu alkaline carbonate complex confirmed the relevant primary and secondary structural properties of syenite alkaline intrusive bodies. The petrography study showed the coarse-grained major and minor minerals such as alkali feldspar, microcline, amphibole, pyroxene, plagioclase, sphene and iron oxides in the samples from PACC. The presence of coarse-grained major and minor minerals in the samples confirmed that all these minerals were formed at deeper depth due to the slow cooling process. Based on the geochemical data, the Pakkanadu alkaline complex highly consisting of alkali elements (Na₂O + K₂O), LILES elements (Sr, Ba, and Rb) and HFSEs (Y, Nb, Zr, Ta, Hf, and Th). The TAS and AFM diagram inferred that PACC mostly be-

Table 2. Trace Elements Composition of Pakkanadu alkaline complex (ppm) at TamilNadu

| Ana-lytete | KAM-26 | KAM-31 | KA-35 | KA-30 | SM-1 | SMP-19 | SMP-16 | SMP-35 | CM-21 | CM-25 | CM-32 | AV-10 | AVA1 | AVA5 | ADA 1 |
|------------|---------|---------|---------|---------|---------|---------|---------|---------|---------|---------|---------|---------|---------|---------|----------|
| Sc | 2.70 | 6.68 | 0.47 | 1.28 | 5.64 | 6.80 | 0.95 | 0.54 | 5.81 | 0.49 | 0.51 | 0.71 | 3.91 | 3.89 | 1.94 |
| V | 114.01 | 243.71 | 73.66 | 243.74 | 42.64 | 42.45 | 49.42 | 55.63 | 95.02 | 101.53 | 71.94 | 63.91 | 115.36 | 131.71 | 9.12 |
| Cr | 39.65 | 10.49 | 12.90 | 9.00 | 9.65 | 7.84 | 6.47 | 13.71 | 9.91 | 11.84 | 8.93 | 12.80 | 29.47 | 31.34 | 5.08 |
| Co | 8.37 | 15.01 | 5.52 | 14.91 | 2.41 | 2.04 | 2.32 | 16.43 | 4.65 | 5.02 | 3.02 | 7.66 | 14.68 | 14.97 | 1.62 |
| Ni | 24.39 | 21.93 | 12.31 | 9.69 | 11.38 | 14.60 | 10.01 | 17.77 | 13.55 | 11.19 | 10.44 | 11.18 | 28.00 | 32.90 | 8.22 |
| Cu | 9.63 | 18.90 | 21.90 | 19.74 | 15.69 | 16.39 | 13.21 | 20.93 | 17.89 | 15.50 | 15.34 | 16.07 | 9.16 | 11.68 | 6.40 |
| Zn | 27.97 | 136.21 | 78.29 | 197.87 | 210.35 | 241.37 | 180.53 | 287.57 | 209.90 | 261.08 | 237.74 | 92.61 | 34.04 | 32.58 | 15.06 |
| Ga | 17.72 | 7.50 | 5.02 | 7.98 | 6.25 | 6.45 | 6.38 | 8.28 | 5.52 | 6.63 | 6.10 | 5.05 | 17.54 | 19.64 | 14.02 |
| Rb | 68.66 | 182.71 | 282.92 | 207.00 | 257.75 | 293.93 | 325.89 | 5.20 | 285.60 | 297.93 | 311.91 | 239.42 | 78.23 | 71.41 | 245.49 |
| Sr | 1513.9 | 2490.33 | 2308.70 | 3135.73 | 2259.49 | 2476.50 | 3112.75 | 1357.66 | 1833.94 | 2224.81 | 2024.02 | 1465.58 | 1525.63 | 1374.7 | 1053.05 |
| Y | 3.61 | 40.73 | 10.08 | 42.19 | 16.78 | 18.96 | 22.11 | 6.90 | 12.38 | 16.87 | 8.14 | 6.63 | 3.52 | 4.25 | 6.38 |
| Zr | 77.24 | 63.09 | 24.79 | 67.94 | 52.00 | 90.96 | 96.11 | 26.04 | 50.68 | 105.40 | 32.44 | 16.25 | 72.83 | 47.14 | 80.29 |
| Nb | 4.08 | 8.85 | 3.80 | 9.20 | 5.89 | 2.20 | 1.44 | 1.34 | 4.04 | 5.02 | 3.03 | 1.72 | 2.81 | 2.70 | 1.48 |
| Cs | 0.56 | 0.64 | 1.02 | 0.68 | 0.75 | 1.19 | 1.06 | 0.15 | 1.18 | 1.26 | 1.33 | 1.33 | 0.59 | 0.45 | 0.84 |
| Ba | 2066.98 | 6231.25 | 8043.98 | 8728.86 | 1358.40 | 522.89 | 1852.38 | 1623.78 | 3956.11 | 4487.88 | 4160.34 | 7482.83 | 1487.09 | 2522.88 | 11148.04 |
| Hf | 2.76 | 2.89 | 0.97 | 3.19 | 1.98 | 2.45 | 2.49 | 0.85 | 1.40 | 2.87 | 1.07 | 0.65 | 2.67 | 1.78 | 1.44 |
| Ta | 0.16 | 1.85 | 0.64 | 1.77 | 1.72 | 0.23 | 0.32 | 0.22 | 0.44 | 1.22 | 0.78 | 0.53 | 0.13 | 0.13 | 0.07 |
| Pb | 20.65 | 18.67 | 22.04 | 31.21 | 64.56 | 64.00 | 58.14 | 30.82 | 50.12 | 59.39 | 37.77 | 26.10 | 25.76 | 22.51 | 33.44 |
| Th | 0.63 | 8.78 | 1.28 | 10.39 | 14.05 | 17.79 | 14.63 | 0.47 | 3.97 | 7.79 | 1.08 | 0.71 | 0.67 | 0.33 | 4.02 |
| U | 0.16 | 0.74 | 0.26 | 0.98 | 1.29 | 1.42 | 1.26 | 0.33 | 0.64 | 1.22 | 0.49 | 0.38 | 0.35 | 0.16 | 0.73 |

Table 3. Rare earth elements composition of Pakkanadu alkaline complex (ppm) at Tamil Nadu

| Analyte | KAM26 | KAM-31 | KA-35 | KA-30 | SM-1 | SMP-19 | SMP-16 | SMP-35 | CM-21 | CM-25 | CM-32 | AV-10 | AVA1 | AVA5 | ADA 1 |
|---------|-------|--------|-------|--------|-------|--------|--------|--------|-------|-------|-------|-------|-------|-------|--------|
| La | 14.59 | 66.31 | 13.98 | 74.81 | 33.40 | 28.48 | 31.33 | 18.92 | 28.34 | 37.54 | 10.81 | 4.59 | 8.94 | 4.35 | 15.307 |
| Ce | 26.36 | 133.66 | 35.52 | 150.33 | 63.37 | 58.40 | 60.27 | 33.06 | 51.03 | 58.10 | 24.11 | 13.11 | 22.69 | 10.95 | 23.687 |
| Pr | 2.36 | 18.29 | 5.21 | 20.30 | 7.44 | 7.14 | 7.09 | 3.93 | 6.30 | 8.99 | 3.39 | 2.22 | 2.10 | 1.15 | 2.165 |
| Nd | 8.55 | 69.52 | 19.55 | 75.45 | 23.31 | 22.95 | 22.30 | 13.45 | 22.04 | 32.09 | 12.59 | 9.52 | 8.44 | 4.83 | 7.748 |
| Sm | 1.21 | 13.62 | 3.71 | 14.01 | 3.68 | 3.68 | 3.94 | 2.17 | 3.78 | 5.72 | 2.36 | 2.17 | 1.44 | 0.92 | 1.618 |
| Eu | 0.76 | 4.69 | 1.71 | 4.91 | 1.25 | 1.21 | 1.46 | 1.37 | 1.55 | 2.20 | 1.10 | 1.21 | 0.54 | 0.88 | 1.505 |
| Gd | 0.89 | 10.36 | 2.78 | 10.92 | 2.75 | 2.86 | 3.30 | 1.72 | 2.97 | 4.33 | 1.92 | 1.76 | 1.02 | 0.82 | 1.595 |
| Tb | 0.13 | 1.38 | 0.38 | 1.41 | 0.38 | 0.41 | 0.45 | 0.23 | 0.39 | 0.58 | 0.28 | 0.23 | 0.14 | 0.13 | 0.147 |
| Dy | 0.65 | 7.32 | 1.98 | 7.59 | 2.27 | 2.22 | 2.79 | 1.22 | 2.11 | 3.00 | 1.47 | 1.26 | 0.66 | 0.76 | 0.863 |
| Ho | 0.14 | 1.60 | 0.41 | 1.64 | 0.54 | 0.53 | 0.67 | 0.27 | 0.46 | 0.64 | 0.31 | 0.28 | 0.13 | 0.17 | 0.162 |
| Er | 0.35 | 4.19 | 1.02 | 4.35 | 1.76 | 1.67 | 2.18 | 0.70 | 1.23 | 1.62 | 0.85 | 0.67 | 0.30 | 0.40 | 0.503 |
| Tm | 0.05 | 0.58 | 0.13 | 0.61 | 0.30 | 0.28 | 0.37 | 0.09 | 0.17 | 0.22 | 0.12 | 0.08 | 0.05 | 0.05 | 0.076 |
| Yb | 0.26 | 3.79 | 0.78 | 4.02 | 2.24 | 2.09 | 2.85 | 0.63 | 1.16 | 1.39 | 0.79 | 0.55 | 0.31 | 0.31 | 0.581 |
| Lu | 0.04 | 0.54 | 0.10 | 0.57 | 0.36 | 0.34 | 0.47 | 0.09 | 0.17 | 0.19 | 0.11 | 0.08 | 0.05 | 0.05 | 0.082 |

longs to alkaline nature based on the composition-wise. The Harker variation diagram confirms the syenite intrusion developed by the magmatic differentiation process in PACC region. The higher concentration of trace elements such as enriched LILE (Ba, Sr, and Rb) elements and High Field Strength Elements (HFSEs) indicate that syenite intrusion is derived from the differentiation of magma mantle source. This enriched and depleted pattern indicates that alkaline magma was derived from mantle sources. The moderate Eu anomaly (Eu/Eu*) suggests that the plagioclase accumulation in the syenite bodies comes from at the crystallization process. The samples association in the plot of syn-COLG + VAG and VAG + syn-COLG confirms that the Pakkanadu alkaline carbonate complex is mostly derived from fractionate of mantle-derived magmatic sources.

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Conflict of interest

The authors declare that they have no conflict of interest.

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